Deriving the thermal & compositional structure of Exoplanets: Examples

Prior Talks:

Radiative Transfer CalculationsLessons from Planetary ScienceWhat learned from Exoplanetary Data

This talk:

Consider HD209458b dayside spectra Play with Radiative Transfer models. How do we proceed with an analysis? What do learn?

Main Collaborators: G.Tinetti, M.Swain, P.Deroo

Questions about hot Jupiters

What happens when Jupiter is moved 100 times closer to the Sun and locked into synchronous orbit?

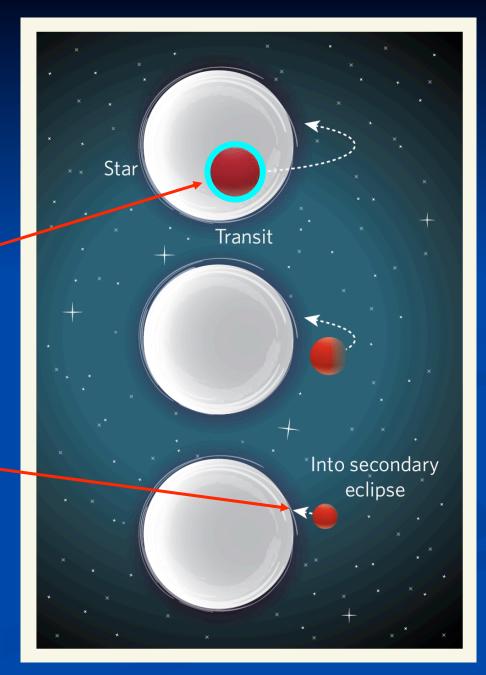
- How efficiently is heat transferred from the day to the night side?
- How does the atmosphere partition energy?
- What governs the chemistry: thermochemistry, photochemistry, ion chemistry?
- What are the dynamics & circulation like?
- What is the ionosphere & magnetosphere like?

Exoplanet Spectroscopy

Transmission during primary eclipse

Emission during secondary eclipse

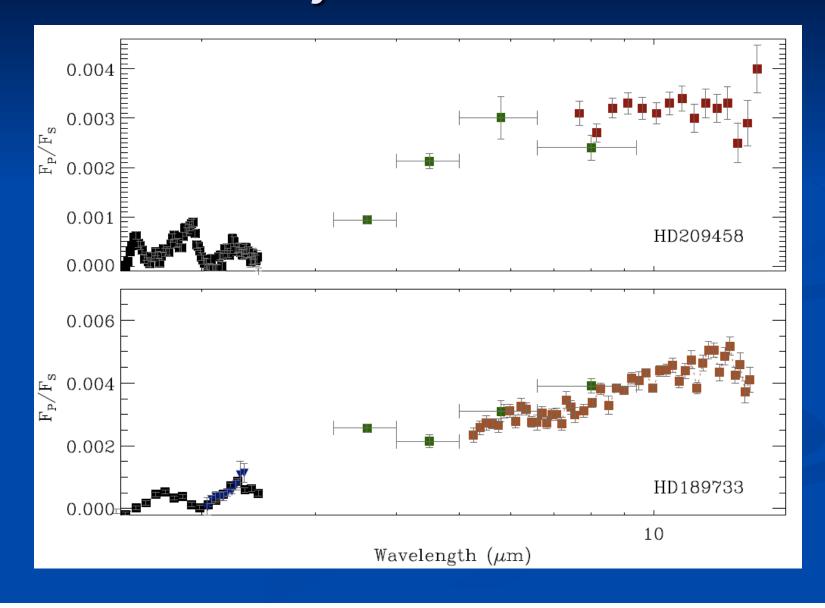
Obital phase variations



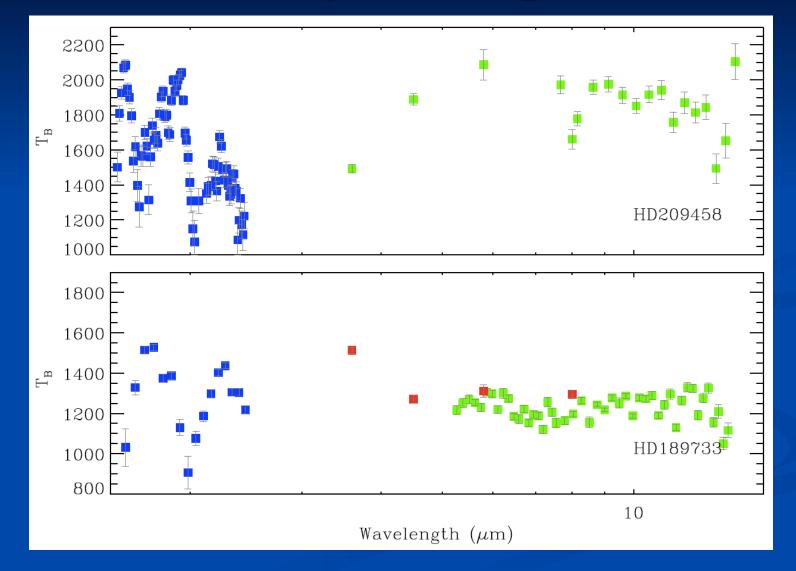
HD209458b

STAR:	
Distance	47 pc
Spectral Type	GOV
Apparent Magnitude V	7.65
Mass	1.01 (± 0.066) Msun
Age	4 (± 2) Gyr
Effective Temperature 5942 K	<u>ref.</u>
Radius	1.146 (± 0.059) Rsun
Metallicity [Fe/H]	0.04
Right Asc. Coord.	22 03 10
Decl. Coord.	+18 53 04
PLANET:	
Mass	0.685 MJ
Semi major axis	0.047
Orbital Period	3.5247 days
Eccentricity	0.07
Radius	1.32 RJ
Inclination	86.677

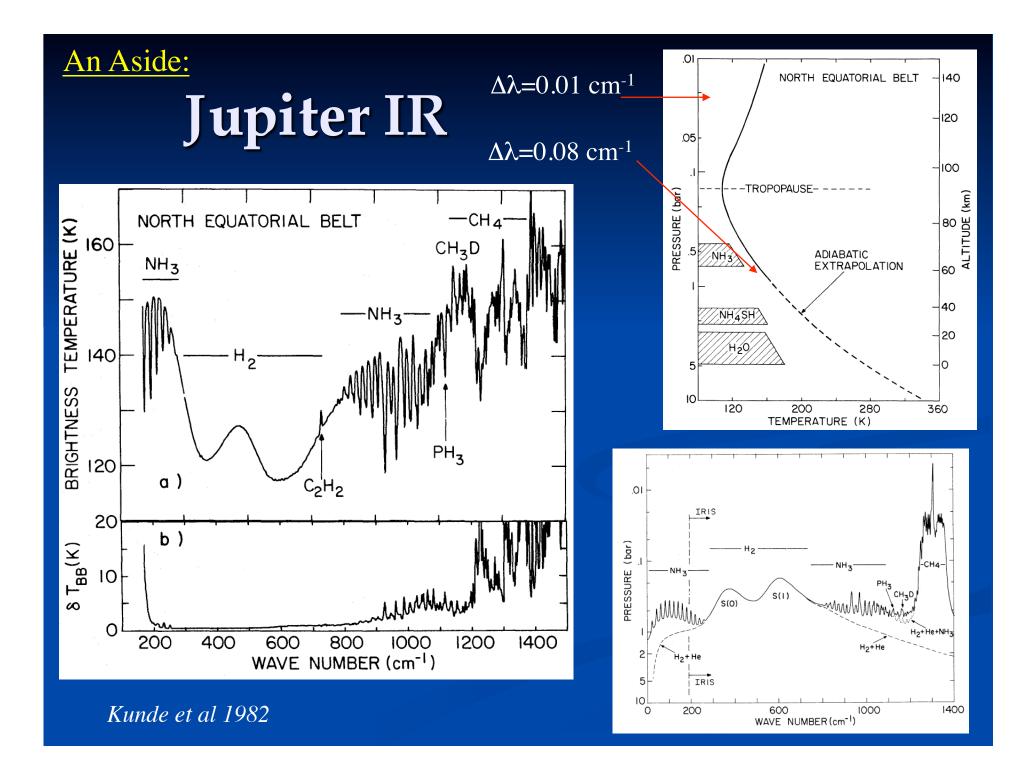
Dayside Flux

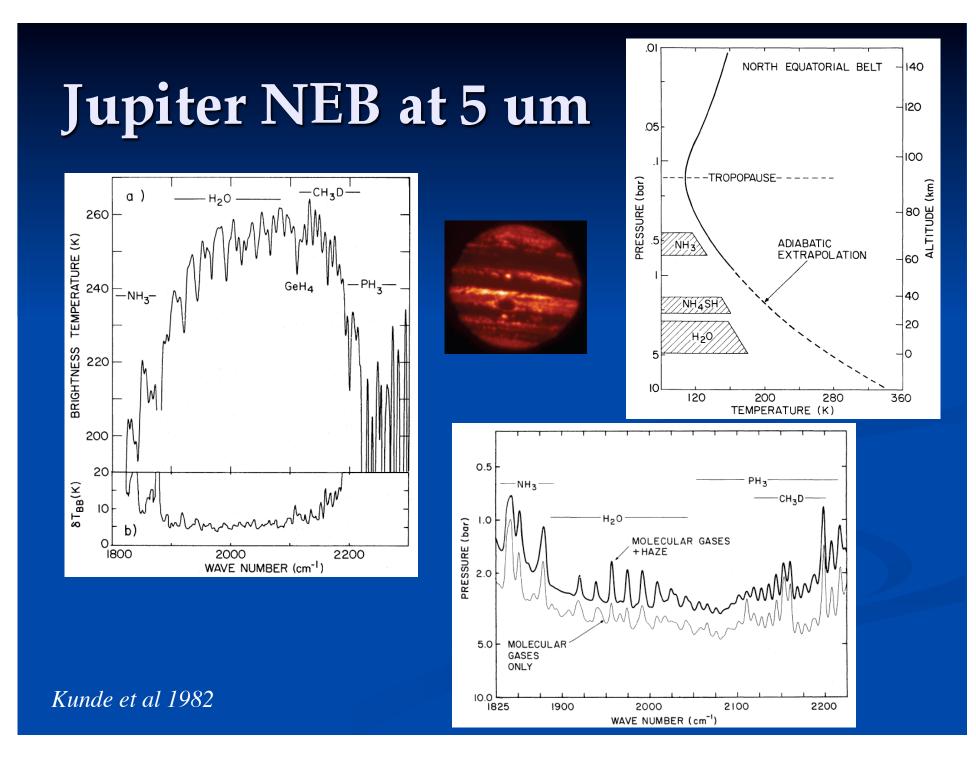


Dayside T_B



Demming et al. 2005, Knutson et al. 2007, Charbonneau et al. 2008, Grillmair et al. 2008, Swain et al. 2009, 2008





Sara's

$$\frac{dI(s,v,t)}{ds} = \varepsilon(s,v,t) - \kappa(s,v,t)I(s,v,t)$$

Radiative Transfer

General Equation:

$$\frac{dI_{\nu}}{d\tau} = -I_{\nu} + S_{\nu}$$

$$S_{v} = \varepsilon/\kappa$$
$$d\tau = \kappa ds$$

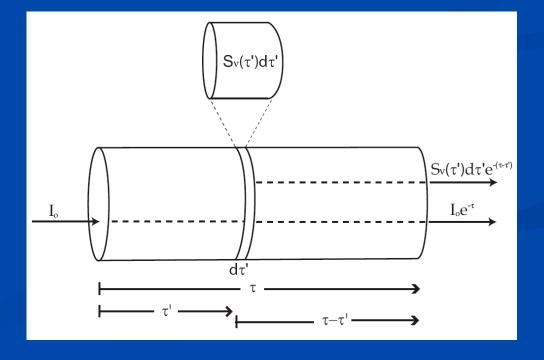
Multiply both sides of equation by e^{τ} .

General Solution:

$$I_{\nu}(\tau_{\nu}) = I_{\nu}(0) \ e^{-\tau_{\nu}} + \int_{0}^{\tau_{\nu}} S_{\nu} e^{-(\tau_{\nu} - \tau_{\nu}')} d\tau'$$

The final intensity equals the original intensity attenuated by absorption & scattering events plus contributions from the source function, S, along the path. The source function includes thermal emission and scattering back into the beam.

 $I_{v} = \text{intensity at wavenumber, } v$ $S_{v} = \text{source fn. at wavenumber, } v$ $\tau_{v} = \text{optical depth at wavenumber, } v$



Radiative Transfer

General Solution:

$$I_{\nu}(\tau_{\nu}) = I_{\nu}(0) \ e^{-\tau_{\nu}} + \int_{0}^{\tau_{\nu}} S_{\nu} e^{-(\tau_{\nu} - \tau_{\nu}')} d\tau'$$

Source terms for scattering of light into the beam make the equation messy, because S then includes the probability, p, that light, $I(\theta', \phi')$, is scattered in the direction of the beam (θ, ϕ)

$$S_{\nu} = \frac{j_{\nu}}{k_{\nu}} = \frac{\kappa_s}{\kappa_e} \int_{4\pi} I_{\nu}(\theta', \phi') \frac{p(\theta', \phi'; \theta, \phi)}{4\pi} d\Omega'.$$

For HD209's IR spectrum no scattering is indicated. We can assume LTE. the source function is the Planck function, B_v , which for a constant local temperature is constant.

$$I_{\nu}(\tau_{\nu}) = I_{\nu}(0) \ e^{-\tau_{\nu}} + B_{\nu}(1 - e^{-\tau_{\nu}})$$

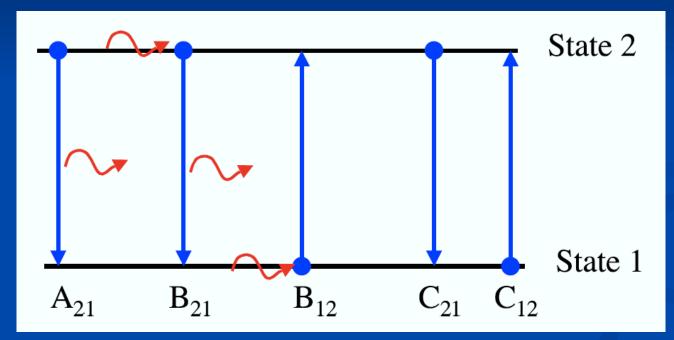
Divide the atmosphere into 80 layers of constant temperature from 10 bars to 10^{-7} bar. Each layer has a specified composition, pressure, temperature, and thus $\tau_v \& B_v$

<u>Radiative</u> <u>Transfer</u> <u>Calculation</u>

$$\begin{array}{c|c} & I_{\nu}^{1} = I_{observed} & \tau_{1} << 1 \\ \hline & T_{1}, P_{1}, \rho_{1}, \Delta \tau_{1} & Z_{2}, \tau_{2} \\ \hline & I_{\nu}^{2} & Z_{2}, \tau_{2} \\ \hline & I_{\nu}^{i} = I^{i+1} e^{-\Delta \tau_{\nu}} + B_{\nu}(T_{i})(1 - e^{-\Delta \tau_{\nu}}) \\ \hline & I_{\nu}^{i} = I^{i+1} e^{-\Delta \tau_{\nu}} + B_{\nu}(T_{i})(1 - e^{-\Delta \tau_{\nu}}) \\ \hline & T_{i}, P_{i}, \rho_{i}, \Delta \tau_{i} & Z_{i}, \tau_{i} \\ \hline & I_{\nu}^{i+1} & Z_{i+1}, \tau_{i+1} \\ \hline & I_{\nu}^{i} = B_{\nu}(T_{N}) & Z_{N}, \tau_{N} >> 1 \\ \hline & T_{N} & \text{or surface} \end{array}$$

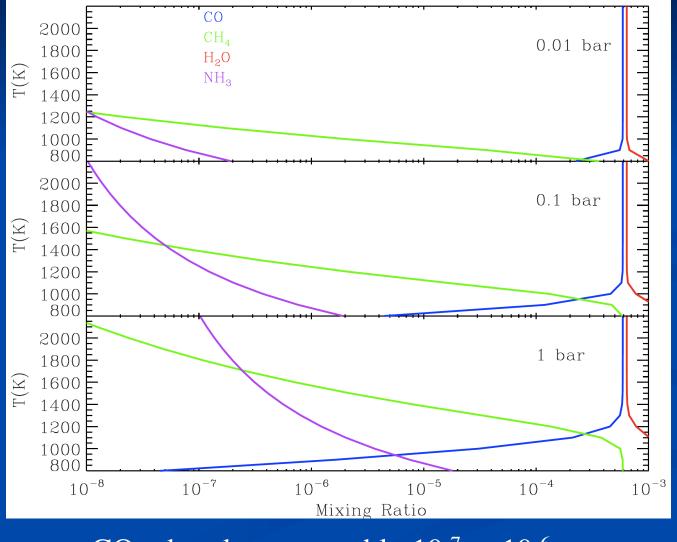
Range of Brightness Temperatures 800 K < T_B <2200 K

If LTE, then T_B is a measure of the probed temperatures



Hypothetical molecule w. 2 discrete energy levels changes states by spontaneous & stimulated emission, absorption, and collisional excitation & de-excitation. If collisions dominate, we have LTE.

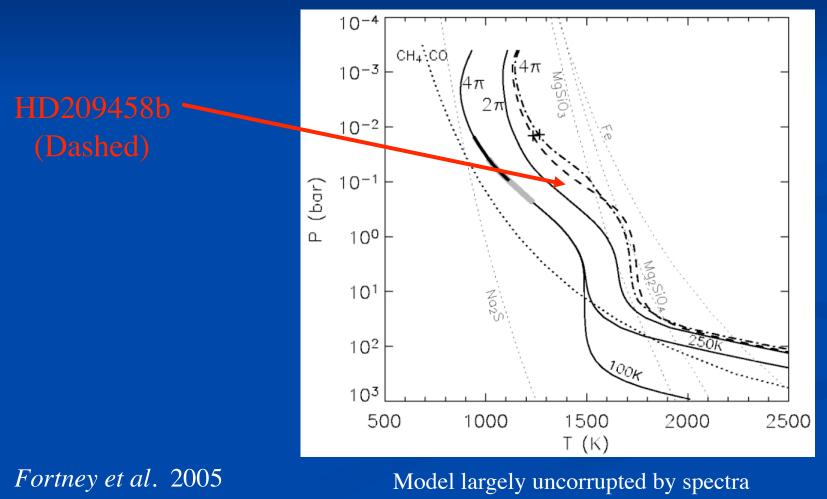
Thermochemical Equilibrium Composition



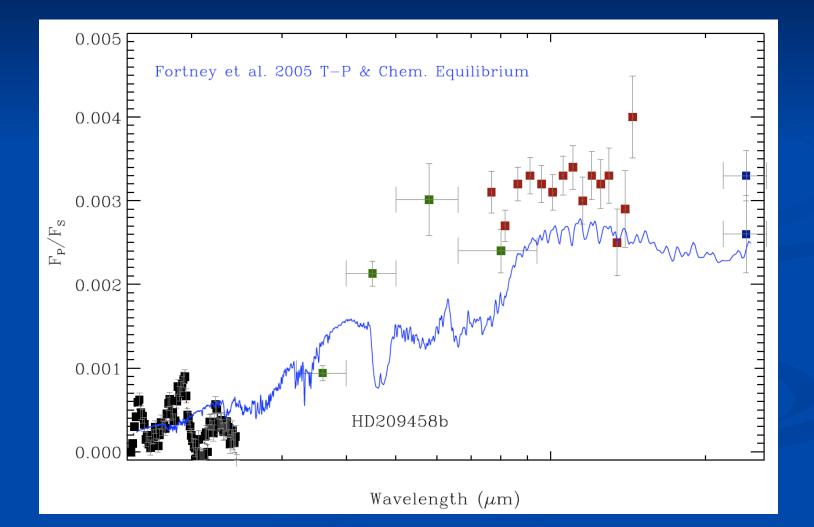
 CO_2 abundance: roughly 10⁻⁷ to 10⁻⁶

Estimate Temperature Profile

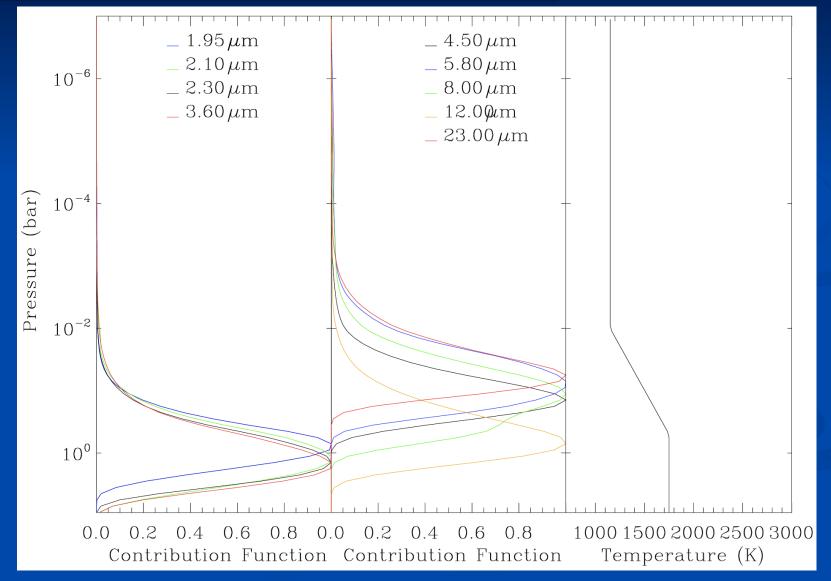
Assume thermochemical equilibrium abundance with solar metalicity Heat fully distributed from day to night (4π reradiated heat)



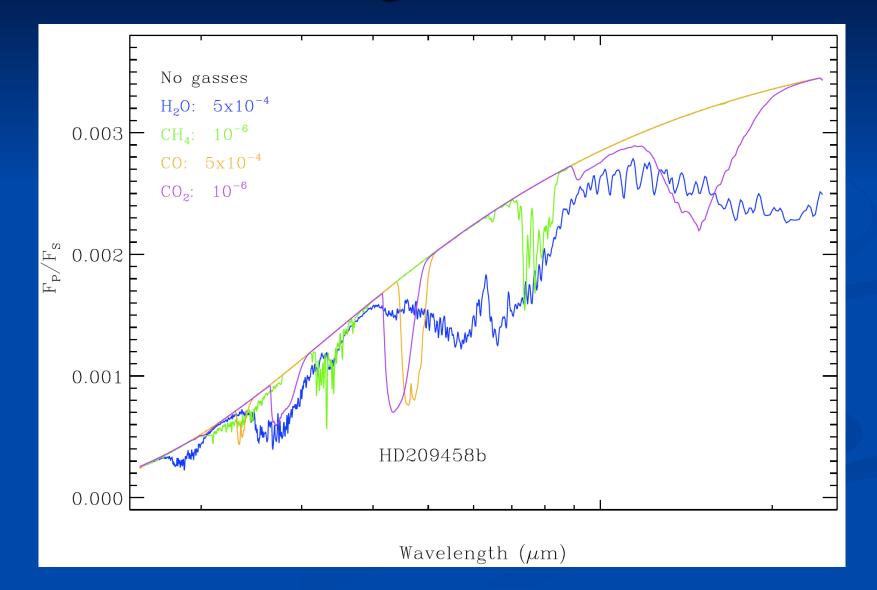
Fit the HD209458b dayside data?



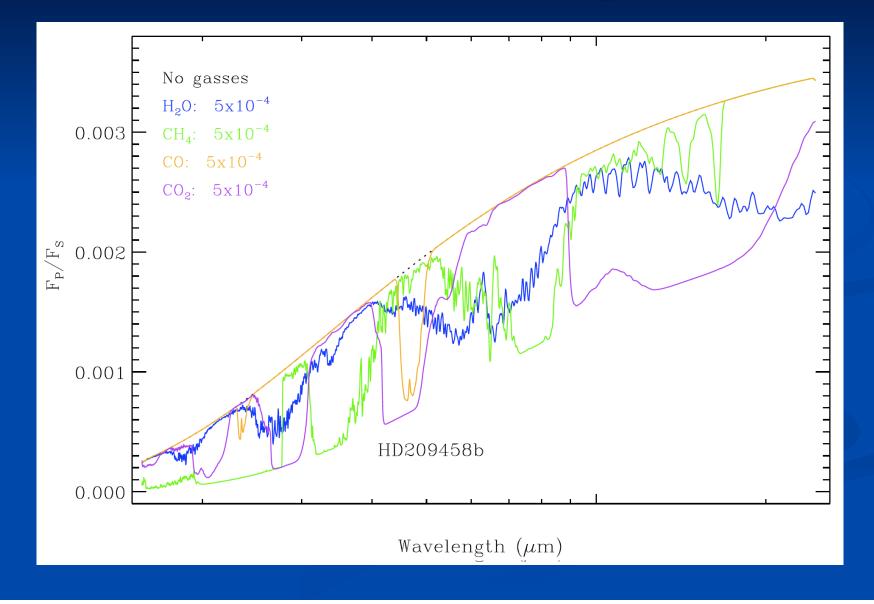
Contribution Functions



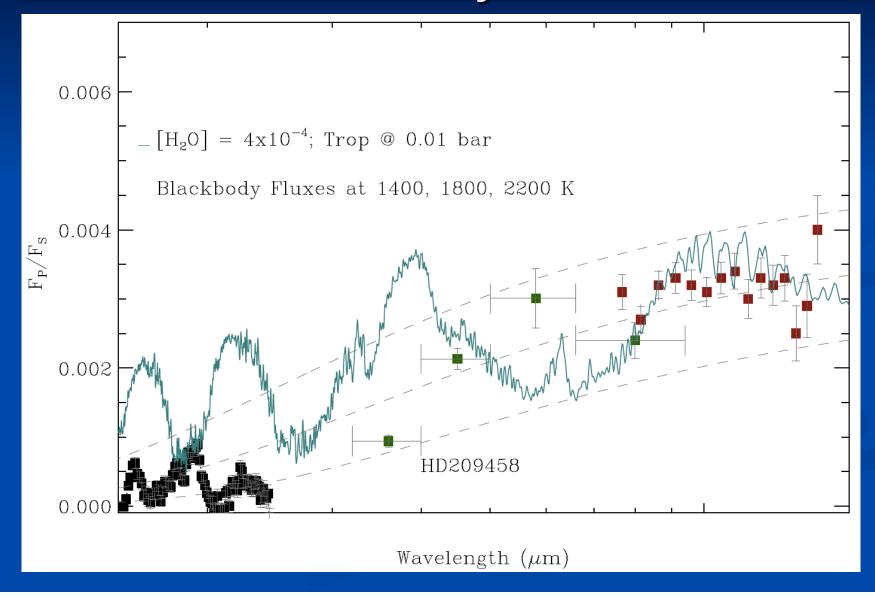
Affects of each gas: water dominates



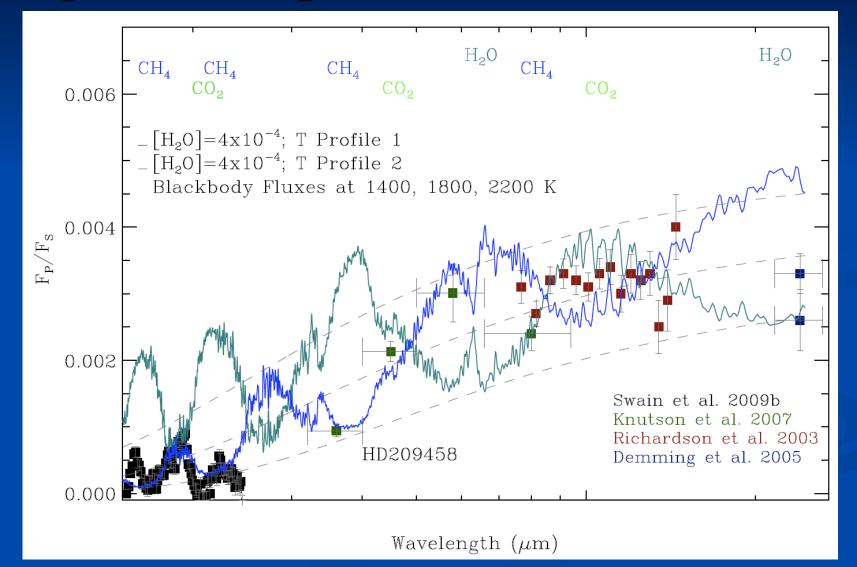
Similar abundances: CH₄ dominates



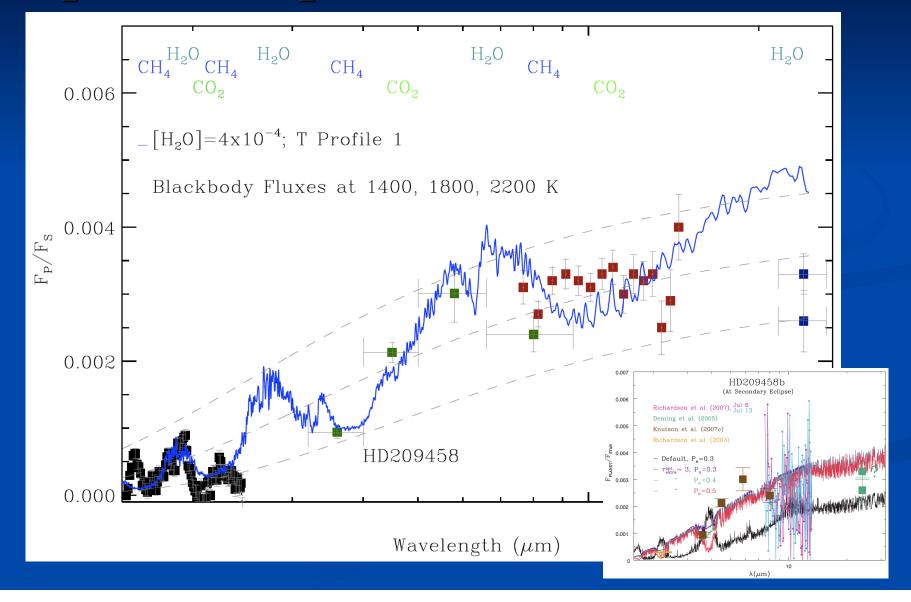
Consider water only & no inversion



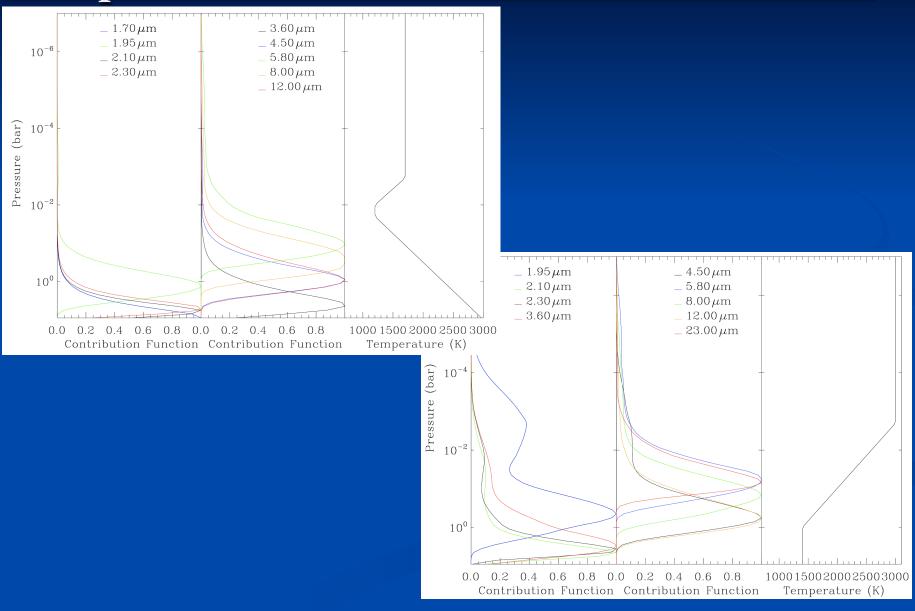
Flip the Temperature



Flip the Temperature



Temperatures & Contribution Functions



Contribution functions

Emerging energy, assuming LTE & no incident radiation:

$$I_{\nu}(\tau_{\nu}) = B_{\nu}(\tau_{surf}) \ e^{-\tau/\mu} + \frac{1}{\mu} \int_{\tau_{surf}}^{0} B_{\nu} e^{-\tau/\mu} d\tau$$

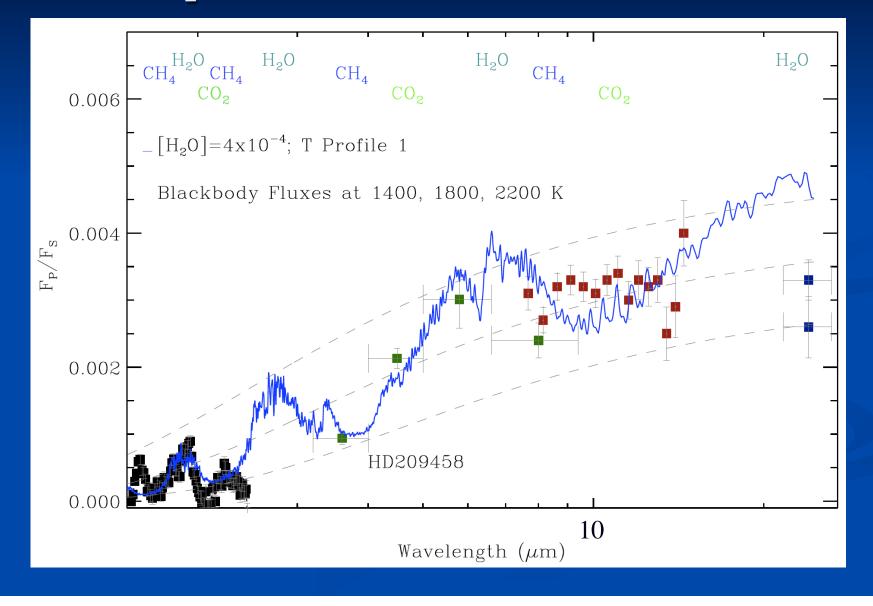
For a bottomless atmosphere:

$$I_{\nu}(\tau_{\nu}) = \frac{1}{\mu} \int_{\infty}^{0} B_{\nu} e^{-\tau/\mu} d\tau$$

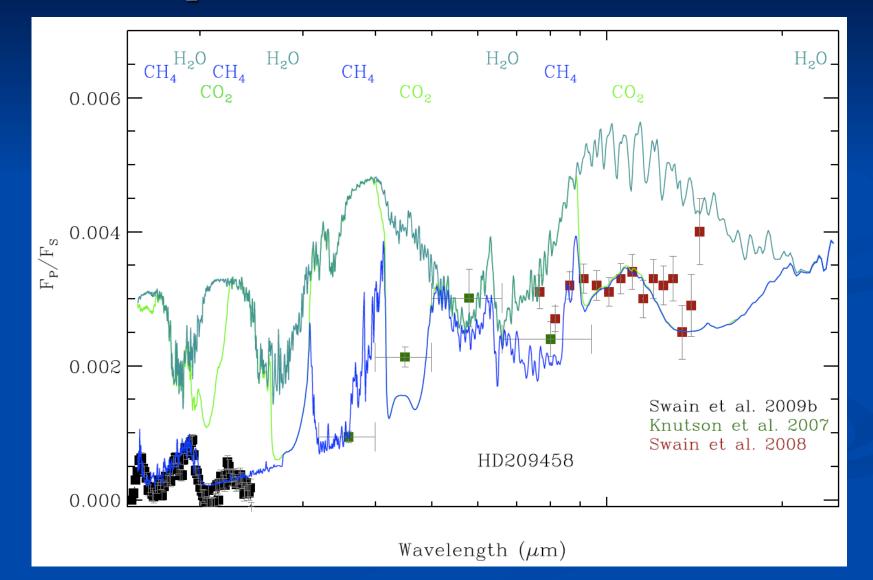
$$I_{\nu}(\tau_{\nu}) = \int_{\infty}^{0} CF(P) \ d\ln P$$

$$CF(P) = B_{\nu}e^{-\tau/\mu}\frac{d(\tau/\mu)}{d\ln P}$$

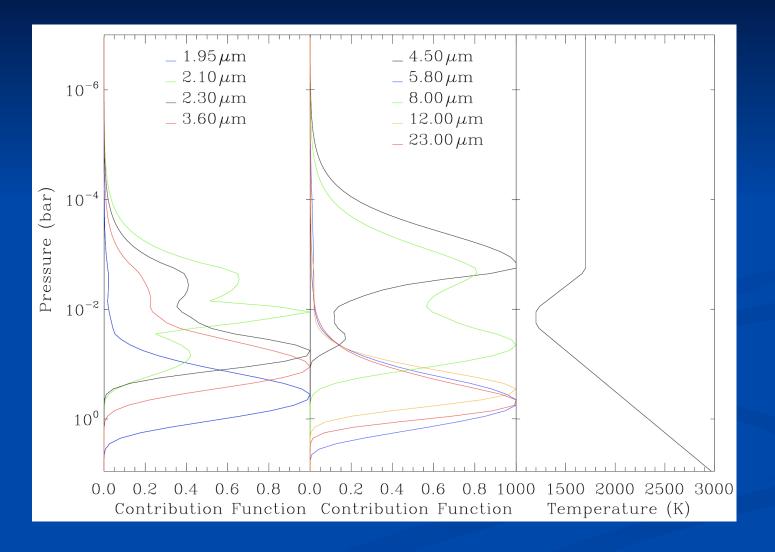
Let CH₄ define the 2 um features



Let CH₄ define the 2 um features



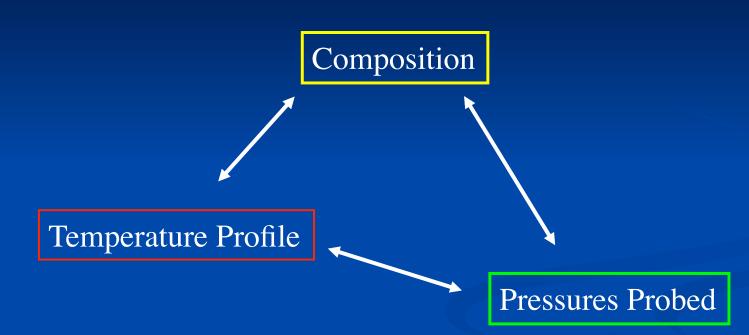
Temperature Inversion indicated by 2 points!



Is this the final solution?

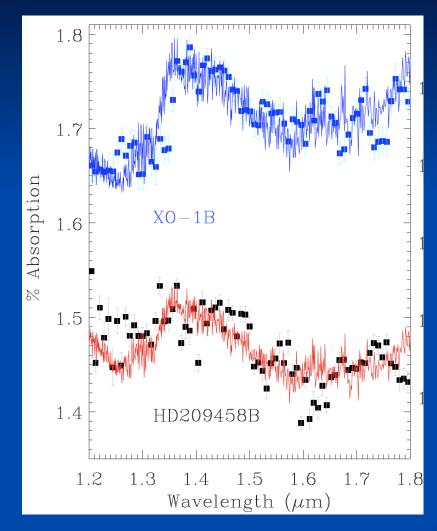


Situation: Leigh's Lesson # 7



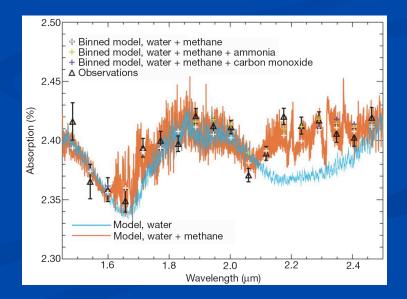
Other constraint: Energy out = Energy in

Limb Transmission



(Tinetti et al. in preparation)

Spectra are consistent! Can't discuss analysis now.



HD189733b (Swain et al. 2008)

Aside to Jupiter

Occultation of β **Scorpii by Jupiter**

(2)

(3)

Consider the ideal gas law,

$$\rho = \frac{Pm}{RT}$$

and integrate this expression:

$$\int_{P_0}^{P} \frac{dP}{P} = -\int_{z_0}^{z} \frac{gm}{RT} dz$$

That is:

$$P = P_0 \ exp[-\int_{z_0}^{z} \frac{gm}{RT} dz] = P_0 \ exp[-\int_{z_0}^{z} \frac{dz}{H}].$$

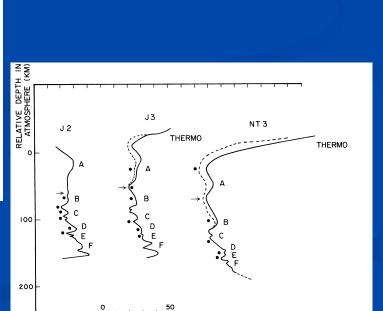
Here we have introduced the pressure scale height, H:

$$H = \frac{RT}{mg}$$

Note that for an isothermal atmosphere (over little change in g), H is constant and is the e-folding distance for changes in ρ and P. That is:

 $P = P_0 \ exp[-z/H].$

-> Jupiter has a hot high atmosphere (that is T > 300 K)



50 IOO REFRACTIVITY SCALE HEIGHT (KM)

Hubbard et al. 1972

Particulates?

0.2-1.0 um Haze particles

cloudless

1) Effects of particulates.

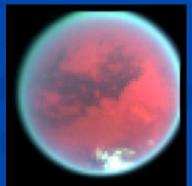
- No indication of its effects in mid-IR
- Indications at optical wavelengths



Titan @ ~0.6 um

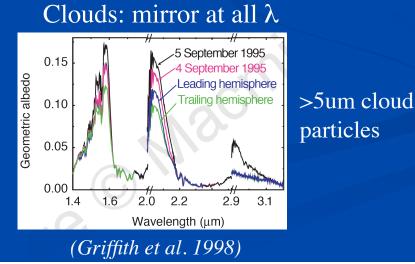


Titan @ 2-5 um



2.50 2.40 2.40 2.30 2.30 2.30 4.57/NICMOS SHIZER 2.30 4.57/NICMOS SHIZER 10 Courtesy F. Pont

Haze: effects short λ



Assume a thermal profile?

1) Radiative Equilibrium?

Conserve flux* throughout the atmosphere F_{in} F_{out}

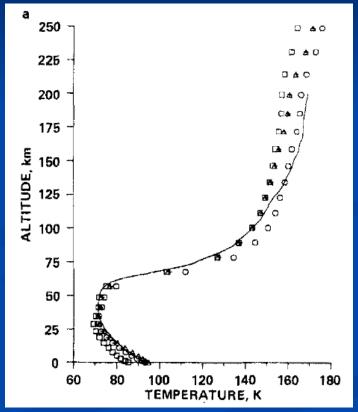
$$\frac{dT}{dt} = -\frac{1}{\rho c_p} \frac{dF}{dz}$$

2) GCM derived profile?

- Continual dialog.
- Clearly distinguish observational from theoretical constraints

• Flux is the integrated intensity component in a particular direction and over all frequencies. It represents the energy / time / area propogated in a particular direction.

Titan Structure



Radiative equilibrium on Titan ($T_{eff} = 85 \text{ K}, \tau_{rad} = 148 \text{ yrs}$)

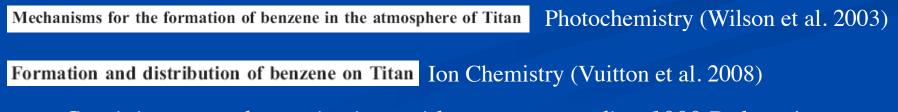
Should assume the chemistry?

Thermochemical Equilibrium?
 Photochemical Model?

Both too early perhaps Again there is a dialog

after

Now need to consider $C_6H_6 = 1-7 \times 10^{-10}$ at 1 mbar 10⁻⁵ at 1 µbar



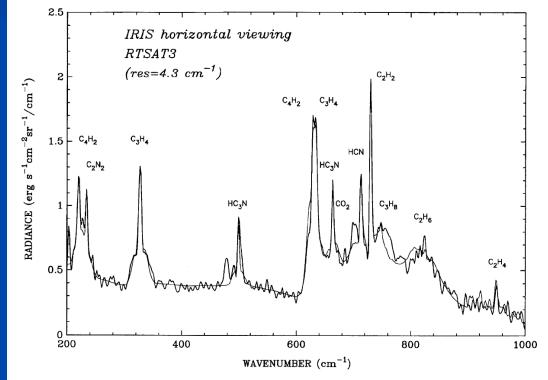
Cassini measured negative ions with masses exceeding 1000 Daltons!

Ignored minor constituents

1) A problem? Yes, but...

- Need features (data)
- Little info on hot lines*

Minor species dominate Titan's spectrum!



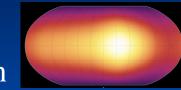
Titan IR spectrum

* We have hot lines information on H_2O , CO, CO_2 , and some CH_4

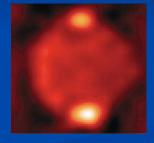
Ignored Stuff

1) Longitudinal variations

- Increasingly more information
- 2) Latitudinal variations
 - No data yet

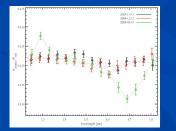


HD189733b



Jupiter @ 3.3 umAurora heated H_3^+ emissions

- 3) Temporal variations
 - Mysterious data under analysis



Deroo in prep

Wish List

1) Coincident observations over large wavelength regions

- Observed variations make modelers nervous
- 2) Fuller spectral coverage data
 - This will refine models (of the largely featureless spectra)
- 3) High resolution spectra
 - Resolve temperature through line strengths
- 4) Spectra of more objects
 - Study processes in a range of conditions
- 5) Line parameters for hot transitions (esp. CH_4)
 - Necessary



Outline

1) Consider an analysis of HD209458b's dayside spectra

- Estimate temperatures, composition and play with RT
 - What assumptions can we make?
 - How well can we constrain this beast?
 - What are the challenges?
- 2) Consider, episodically, jovian or planetary counterparts
 - Understanding composition, temperatures, and structure
- 3) What do we need in the future to understand exoplanets?
 - Additional Observations?
 - Additional Theoretical Calculations?
 - Additional Lab work



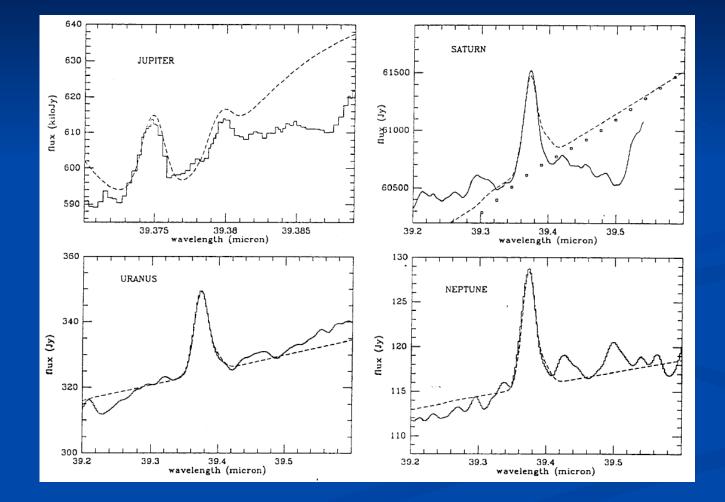
Jupiter (Case Example)

How does one determine the vertical profiles of:

temperature $[H_2]$ [He] $[CH_4]$ $[NH_3]$ $[H_2O]$ [CO] $[CO_2]$

H2O 2-5 bar in Jupiter

Stratospheric water

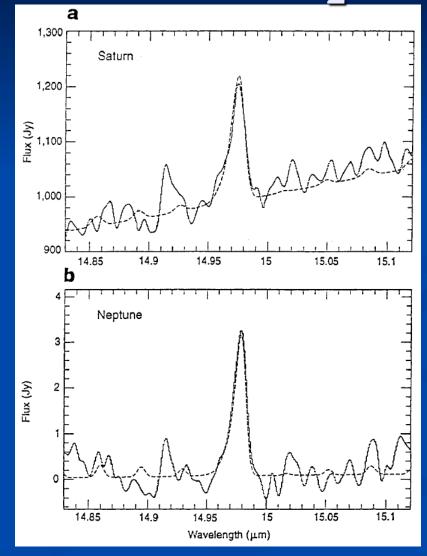


 $[H_2O]=10^{-9}$ at roughly 10 mbar - 10 ubar

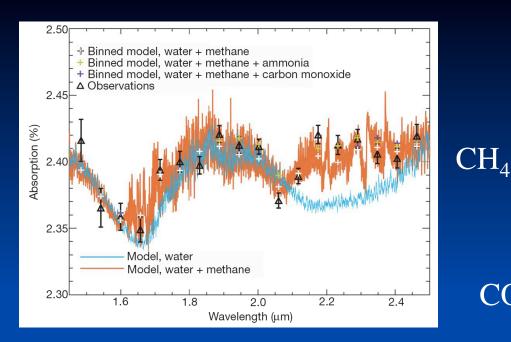
Detection of CO₂

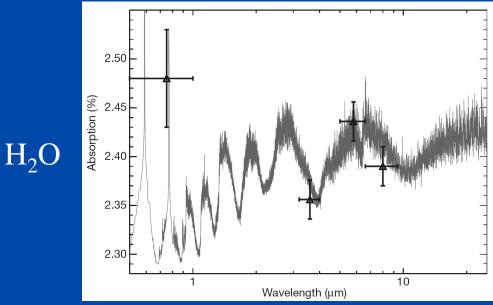
De Graauw et al. 1997

Feuchtgruber et al. 1997

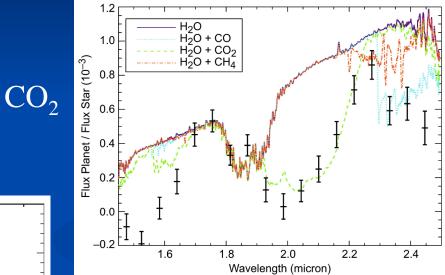


"Dieu n'est pas pour les gros bataillons, mais pour ceux qui tirent le mieux" ("God is not on the side of the big battalions, but of the best shots") – Voltaire



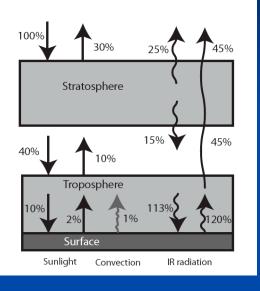


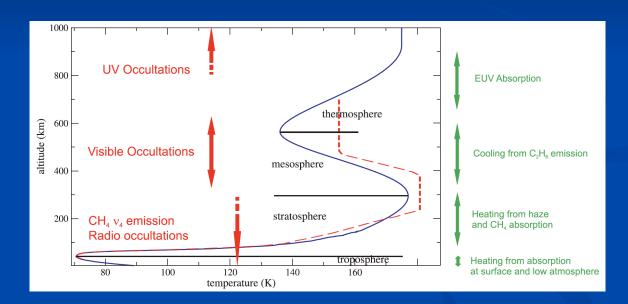
Giovanna & gang



How is heat partitioned in the atmosphere?

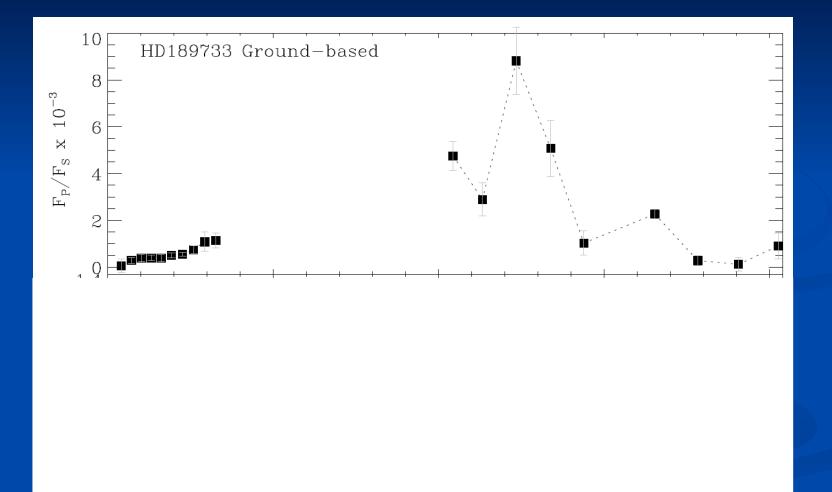
TITAN





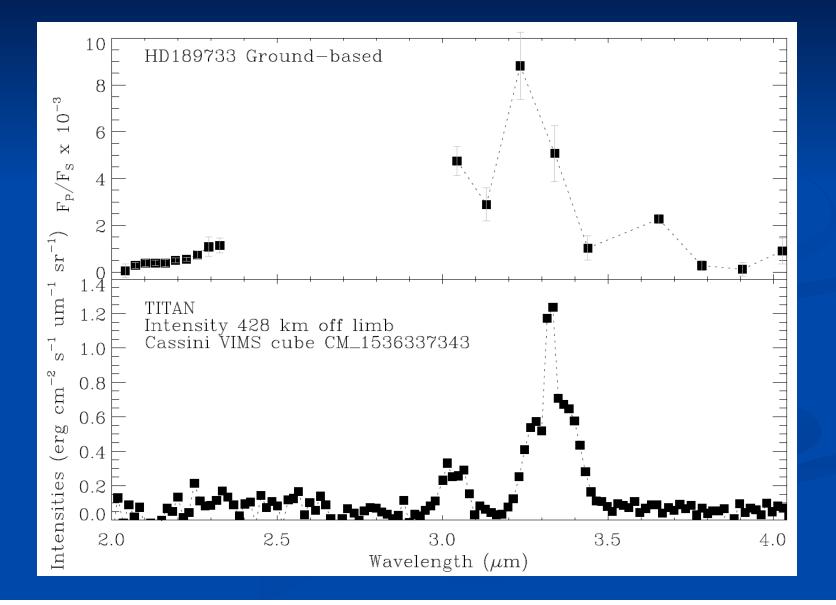
Calculate the radiative balance of the atmosphere. Compare to observations to constrain dynamical effects.

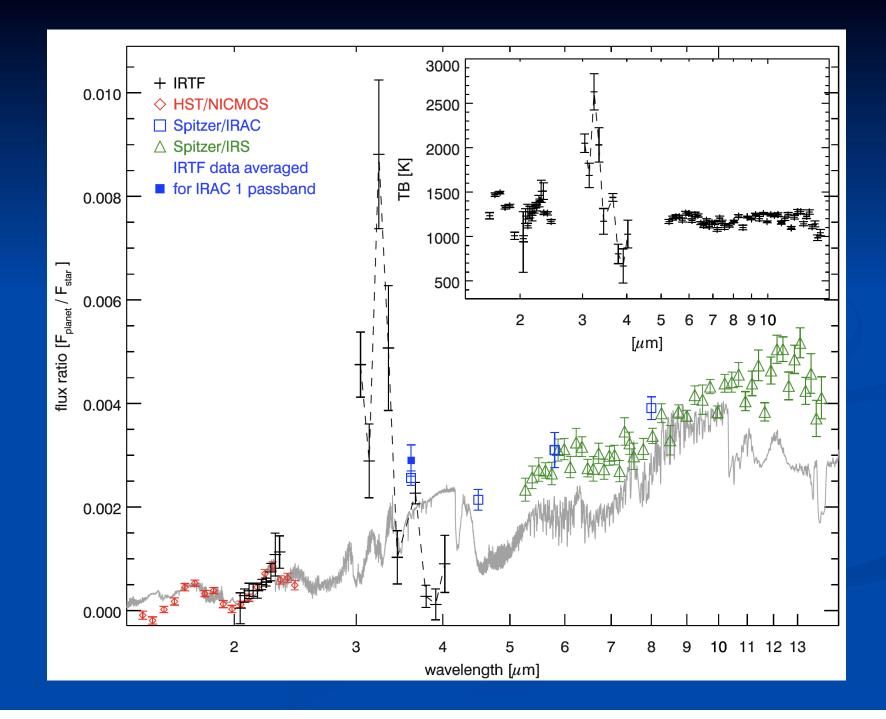
New Result



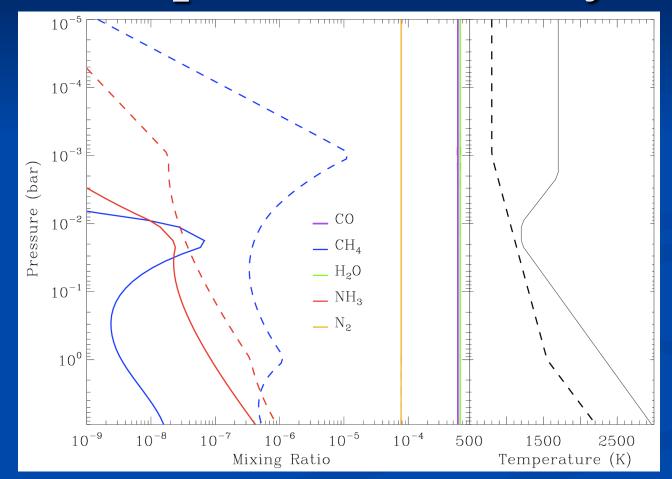
U (1 /

New Result





What determines the atmospheric chemistry?



Thermochemical equilibrium calculations are not enough

How do planetary processes scale with a planet's mass and semi-major axis?

Exoplanet	Mass	Semi-major axis	Eccentricity	Primary	${f Spectra}$
HD209459B	$0.63 M_J$	$0.0467 \ AU$	0.07	${ m G0V} \ / \ { m 5942} \ { m K}$	UV to 24 μm
HD189733B	$1.13 M_J$	$0.0310 \ AU$	0	${ m K2V}$ / 4980 K	0.5-24 $\mu \mathrm{m}$
XO-1B	$0.90 M_J$	$0.0488 \ AU$	0	G1V	1.4-2.4 $\mu \mathrm{m}$
TrES-1	$0.61 \ M_J$	$0.0393 \; AU$	0	$\mathrm{K0V}$	photometry
HD80606B	$4.00 M_J$	0.453 AU	0.936	G5 / 5370 K	photometry
Gl436B	$0.072 \ M_J$	$0.0287 \ AU$	0.15	M2.5 / 3684 K	photometry
Gl876D	$0.018 \ M_J$	$0.0208 \ AU$	0	M4V / 3350 K	photometry

